Imaging Seismic Attenuation at the Brady Geothermal Field Using Interferometry

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The PoroTomo "Natural Lab"

1500-by-500 meter natural laboratory at the Brady EGS field



geologic obstacles by Coolbaugh

Key goal: understand how fluids travel from shallow aquifers, through faults and fractures, to deep geothermal reservoirs.

Seismic, geodetic, and hydraulic technologies are applied to fully characterize the rock mechanical properties.

Conceptual model: Highly permeable conduits along faults channel fluids from shallow aquifers to the deep geothermal reservoir tapped by the production wells.



A large seismic array was deployed over the geothermal field.



geophones (yellow in lab, cyan external to lab)

238 nodes recorded more than two weeks of continuous data

 local and regional earthquakes, vibroseis sweeps, local traffic noise, and the ambient seismic wavefield.

"Virtual Earthquakes": sources of seismic energy



vibroseis points (green). Highway I-80 parallel to natural lab. geologic obstacles from Coolbaugh, faults from Faulds Types of data: ambient noise, coda, active sources

- Focus structure between seismometers
- Simple estimate of the GF
- Perfect location and timing constraints



Using Vibroseis Sweeps as a source of coherent energy



vibroseis points (green), geophones (yellow) geologic obstacles from Coolbaugh, faults from Faulds

Vibroseis sweeps at points surrounding the site during the experiment.

Only short records of the active sweeps are required to estimate high frequency GFs (up to 30 Hz).

Computationally much less expensive than ANC.



We see large lateral heterogeneity in seismic velocity



surface expression of faults in white, fumaroles (orange circles)

geologic obstacles from Coolbaugh, faults from Faulds

- Inverted for Vp, Vs, Qs, Qp
- At the surface Vp varies by more than 50%.
- Anomalies align with mapped hotspots and faults.

Vp/Vs contrasts appear to map subsurface fabric



Vp/Vs (50 m)

50 m depth of faults in white, fumaroles (orange circles)

geologic obstacles from Coolbaugh, faults from Faulds

- Inverted for Vp, Vs, Qs, Qp
- Max Vp/Vs decreases rapidly with depth in the top 100 m.
- Surface Vp/Vs varies between 3.5-7.

Attenuation of seismic energy increases in regions that are hot or heavily fractured.



Qs (150 m)

150 m depth of faults in white, fumaroles (orange circles)

geologic obstacles from Coolbaugh, faults from Faulds

- Inverted for Vp, Vs, Qs, Qp
- Each has different sensitivity to material properties (temperature, porosity, fluid content and composition.)

Qs and Qp have different sensitivity to fluid saturation



Qs/Qp anomalies align on faults



Pressures were changed in four stages over the two week experiment.



operations at recorded in three monitoring boreholes (Feigl, 2017)

Phase arrival times are nearly identical, but amplitudes changed measurably after site shutdown



Changes in amplitude are concentrated in fault bounded blocks.





Static image of attenuation during normal operations (Stage 1)

Dynamic changes in seismic amplitudes as operations changed

- blue: more efficient propagation
- red: more attenuated

Conclusions

PoroTomo objective:

- Characterize spatial distribution and monitor temporal changes rock mechanical properties of EGS reservoir in 3D.
- Infer Young's modulus, Poisson's ratio, saturation, porosity density. (ref. Kurt Feigl SGW 2017).
- Using interferometry we obtain images of Vp, Vs, Qp and Qs at high resolution.
- Brady geothermal field is highly heterogeneous in the top 100 m.
- Vp/Vs anomalies largely follow fault structure.
- Qs/Qp is largest following major faults and to greatest depth beneath the injection point and near the fumaroles.
- Changing pressures appear immediately in the seismic amplitudes, concetrated in fault bounded blocks.



Poster presentation SSA:

Seismic Tomography at Brady Geothermal Field, Nevada with Dense Nodal and Fiber-Optic Seismic Arrays, Nayak et al.

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Reinisch et al. in prep

Comparison with conceptual model

