

Robustness of κ_0 measurement: insight from a site in the low-to-moderate seismicity context of southeastern France

Presented by
Vincent PERRON

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ETH

UNIVERSITÉ

Grenoble

Alpes

Eidgenössische Technische Hochschule Zürich Swiss Federal Institute of Technology Zurich

Acceleration FAS Individual κ_r κ_r represent the attenuation along the $e^{-\pi \kappa_r f}$ Record path between the For a given earthquake located at epicentral distance R_e , κ_r is define by the relation: source and the $A(f) = A_0 e^{-\pi \kappa_r f}, \quad f_1 < f < f_2$ recording site fmax Acceleration (t, f) Source No influence of ω^{-2} model the source f 2ini 10-2 4 Hz Site Main assumptions f_c Frequency $e^{-\pi \kappa_r f}$ ttenuation independent Acceleration FAS (m.s⁻¹.Hz^{-1/2}) PathPath attenuation Source parameter S-waves Noise R_e Adapted from Berge-Negligible $\overset{\mathfrak{H}}{\overset{\mathfrak{g}}{\overset{\mathfrak{g}}{\overset{\mathfrak{g}}{\overset{\mathfrak{g}}{\overset{\mathfrak{g}}{\overset{\mathfrak{g}}{\overset{\mathfrak{g}}{\overset{\mathfrak{g}}{\overset{\mathfrak{g}}{\overset{\mathfrak{g}}{\overset{\mathfrak{g}}{\overset{\mathfrak{g}}{\overset{\mathfrak{g}}{\overset{\mathfrak{g}}{\overset{\mathfrak{g}}{\overset{\mathfrak{g}}{\overset{\mathfrak{g}}{\overset{\mathfrak{g}}{\overset{\mathfrak{g}}{\overset{\mathfrak{g}}{\overset{\mathfrak{g}}{\overset{\mathfrak{g}}{\overset{\mathfrak{g}}{\overset{\mathfrak{g}}{\overset{\mathfrak{g}}{\overset{\mathfrak{g}}{\overset{\mathfrak{g}}{\overset{\mathfrak{g}}{\overset{\mathfrak{g}}{\overset{\mathfrak{g}}{\overset{\mathfrak{g}}{\overset{\mathfrak{g}}{\overset{\mathfrak{g}}{\overset{\mathfrak{g}}{\overset{\mathfrak{g}}{\overset{\mathfrak{g}}{\overset{\mathfrak{g}}{\overset{\mathfrak{g}}{\overset{\mathfrak{g}}{\overset{\mathfrak{g}}{\overset{\mathfrak{g}}{\overset{\mathfrak{g}}{\overset{\mathfrak{g}}{\overset{\mathfrak{g}}{\overset{\mathfrak{g}}{\overset{\mathfrak{g}}{\overset{\mathfrak{g}}{\overset{\mathfrak{g}}{\overset{\mathfrak{g}}{\overset{\mathfrak{g}}{\overset{\mathfrak{g}}{\overset{\mathfrak{g}}{\overset{\mathfrak{g}}{\overset{\mathfrak{g}}{\overset{\mathfrak{g}}{\overset{\mathfrak{g}}{\overset{\mathfrak{g}}{\overset{\mathfrak{g}}{\overset{\mathfrak{g}}{\overset{\mathfrak{g}}{\overset{\mathfrak{g}}{\overset{\mathfrak{g}}{\overset{\mathfrak{g}}{\overset{\mathfrak{g}}{\overset{\mathfrak{g}}{\overset{\mathfrak{g}}{\overset{\mathfrak{g}}{\overset{\mathfrak{g}}{\overset{\mathfrak{g}}{\overset{\mathfrak{g}}{\overset{\mathfrak{g}}{\overset{\mathfrak{g}}{\overset{\mathfrak{g}}{\overset{\mathfrak{g}}}{\overset{\mathfrak{g}}{\overset{\mathfrak{g}}{\overset{\mathfrak{g}}}{\overset{\mathfrak{g}}{\overset{\mathfrak{g}}}{\overset{\mathfrak{g}}{\overset{\mathfrak{g}}}{\overset{\mathfrak{g}}{\overset{\mathfrak{g}}}{\overset{\mathfrak{g}}{\overset{\mathfrak{g}}}{\overset{\mathfrak{g}}}{\overset{\mathfrak{g}}{\overset{\mathfrak{g}}}{\overset{\mathfrak{g}}}{\overset{\mathfrak{g}}}{\overset{\mathfrak{g}}}{\overset{\mathfrak{g}}}{\overset{\mathfrak{g}}}{\overset{\mathfrak{g}}}{\overset{\mathfrak{g}}}{\overset{\mathfrak{g}}}{\overset{\mathfrak{g}}}{\overset{\mathfrak{g}}}{\overset{\mathfrak{g}}}{\overset{\mathfrak{g}}}{\overset{\mathfrak{g}}}{\overset{\mathfrak{g}}}{\overset{\mathfrak{g}}}{\overset{\mathfrak{g}}}{\overset{\mathfrak{g}}}{\overset{\mathfrak{g}}}{\overset{\mathfrak{g}}}}{\overset{\mathfrak{g}}}{\overset{\mathfrak{g}}}}{\overset{\mathfrak{g}}}{\overset{\mathfrak{g}}}{\overset{\mathfrak{g}}}}{\overset{\mathfrak{g}}}{\overset{\mathfrak{g}}}}}}$ ficationThierry et al. (2017) 10⁻⁶ site amplification ^{Sn Id} 10 10 15 20 25 30 35 40 Frequency (Hz) f_0

frequency (Hz)

The site component κ_0

 κ_0 correspond to the site component of κ_r such as:

$$\kappa_r(R_e) = \frac{\kappa_0}{\kappa_0} + \frac{m_\kappa}{\kappa_0} \cdot R_e$$



 $\kappa_r(s)$

Х

Х

Approaches to measure κ_r

κ_{AS} : Anderson and Hough (1984)



- High frequency measurement $> f_c$
- *M_L*>3



κ_{DS} : Biasi and Smith (2001)



Low frequency measurement < f_c
 M_L<1,5

Use of κ_0 :

- Ground Motion Prediction Equations (GMPEs)
- Stochastic modeling
- Host-to-target adjustments
- etc.

II. The study site

The study site

- Seismicity: low-to-moderate
- Recording period: 2^{1/2} years
- Instruments: 7 velocimeters et 2 accelerometers
- About 70 earthquakes usable for κ_r measurements



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• "AS O "DS

10¹

10²

≥ 3

10⁰

Comparison between κ_{AS} and κ_{DS}

 κ_{AS} : Anderson and Hough (1984)

 κ_{DS} : Biasi and Smith (2001)





Frequency-dependence assumption

- The frequency dependence can explain a parts of the discrepancy between m_{k_AS} and m_{k_DS}
- No quantifiable effect between κ_{0_AS} and κ_{0_DS}



■ Negligible amplification influence assumption



- Non-negligible influence on *κ*₀ at rock sites
- Very probable and unpredictable influence on κ_0 at soil sites



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231 inverted Fourier source

spectra from the L'Aquila sequence



V. Conclusions

Conclusions

- Consistent results between κ_{0_AS} and κ_{0_DS} at \longrightarrow rock site, and in agreement with the literature
- For low-to-moderate seismicity areas, the *K_{DS}* approach is a promising alternative to the very uncertain V_{S30}-*K*₀ correlations







- Very probable frequency-dependence of κ
- Noticeable influence of the amplification on the average values of κ_0 , even for rock sites
- Despite *k*₀ is a site parameter, the uncertainty associated to its measurement is clearly and strongly dominated by the source term

Vincent PERRON

vincent.perron@sed.ethz.ch







ETTH Eidgenössische Technische Hochschule Zürich Swiss Federal Institute of Technology Zurich

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